## **PLATES**

# PLATES I-VIII (Figs. 23-74)

Fig. 23 to Fig. 29: The contacts with the underlying and overlying rocks. Fig. 30 to Fig. 38: Miscellaneous. Fig. 39 to Fig. 48: The excavation S of Grandoso. Fig. 49 to Fig. 55: The excavation E of Boñar. Fig. 56 to Fig. 58: Miscellaneous. Fig. 59 to Fig. 74: Thin sections.

Photographs by Mr. W. C. Laurijssen

#### PLATE I

Fig. 23. Contact with the Paleozoic W of Sorribos de Alba; to the left Paleozoic, to the right Voznuevo Formation.

Fig. 24. Contact with the Paleozoic N of La Valcueva. The Voznuevo Formation contains large quartzitic boulders and cobbles of Paleozoic origin (A).

Fig. 25. Contact between the Barrios Formation (Ordovician) and the Voznuevo Formation N of the village of Voznuevo, looking W.

Fig. 26. Hand specimen showing the contact between the Barrios Formation and the Voznuevo Formation. The Voznuevo Formation consists of cross-stratified sand and gravel, immediately overlying the contact.

Fig. 27. Remainders of consolidated Voznuevo material (A) on the unconformity surface N of Voznuevo.

Fig. 28. Cracks in the unconformity surface filled with Voznuevo sand and gravel (arrow) N of Voznuevo.

Fig. 29. Contact with the overlying Vegaquemada Formation SE of Carrocera. Right hand side: Voznuevo Formation; left hand side: Vegaquemada Formation, darker coloured due to a higher content of iron (hydr)oxides.

Fig. 30. Very thinly laminated alternations of sand (dark colour) and clay (light colour) in youngest detailed section of section E. Arrow indicates top.





#### PLATE II

Fig. 31. Strongly burrowed and rooted black clay in Section II. Observer looking upon the plane of stratification.

Fig. 32. Black clay veined by thin sand stringers and burrows, Section J.

Fig. 33. Fluvial fining-upward grading cycles N of Brugos de Fenar (A, B). The upper parts of the units are marked by clay layers (Cl).

Fig. 34. Well-preserved cross-stratification in consolidated sand and gravel, top of Section K.

Fig. 35. Very coarse-grained cross-stratified unit between two clayey horizons (Cl). Note abrupt changes in grain size in a lateral direction. Bottom of Section L.





#### PLATE III

Fig. 36. Very thinly bedded small-scale grading units (from fine-grained sand to clay) in Section III (Section J). Note traceability of even the thinnest laminae. Arrow indicates stratigraphic top.

Fig. 37. Straight to slightly undulating ripple crests exposed on principal bedding plane. Top of Section T.

Fig. 38. Undulating boundary (dashed line) between channels with slightly diverging paleocurrent directions, observer looking downstream. Exposure NE of Boñar.

Fig. 39. Mega cross-stratified unit with backflow in the foresets (BFF) and in the bottomset layer (BFB).

Fig. 40. Detail of backflow in a mega foreset of Fig. 39. Note longitudinal grading in the backflow unit (indicated by arrow).

Fig. 41. Channel with backflow structures (BFB) in bottomset layer.

Fig. 42. Erosive transition between two cross-stratified units with opposed current directions showing large clay balls (Cl); arrows indicate current directions.

Fig. 43. Nose-like intrusion of fine-grained material (N) in mega foreset and backflow cross-stratification (BFB) in bottomset.





#### PLATE IV

Fig. 44. Erosion of a cross-stratified unit (B) by a unit with opposed current direction (A). At the transition clay balls occur (Cl). Arrows indicate current directions.

Fig. 45. Cross-stratified unit, devoid of backflow structures, passing laterally into a cross-stratified unit with distinct backflow (BF).

Fig. 46. Distinct fining-upward grading in mega foresets of a cross-stratified unit.

Fig. 47. Cross-stratified unit with small-scale cross-stratification in the foresets which have the same current direction. 'Coflow in the ripple'.

Fig. 48. Detail of the right hand side of Fig. 47, showing cross-stratification in the mega foreset (MF). Note vertical grading in the mega foreset and longitudinal grading in the 'internal' cross-stratification.

Fig. 49. Cross-stratified unit (a) underlain by finer-grained backflow unit (b) and by coarsergrained bottomset layer (c). For explanation see text.



PLATE V

Fig. 50. Detail of units V to Y (cf. Fig. 52). (BF = Backflow).

Fig. 51. General view of the SW wall of the excavation E of Boñar. Letters and numbers refer to text.





#### PLATE VI

Fig. 52. Detail of the easternmost part of the SW wall of Fig. 51 showing units U, V, W, X, Y and Z. In the upper part of unit U a sudden increase in coarse-grained material can be observed (A).

Fig. 53. Asymmetrically filled channel with supply direction from the E, observer looking NE. Channel filling shows perfectly developed longitudinal grading.

Fig. 54. Mega cross-stratified units with backflow phenomena (BF) in their lower parts (detail of unit Y in Fig. 51).

Fig. 55. Backflow (BF) between almost horizontally layered fine sand to silt layers (detail of unit Y in Fig. 51).

Fig. 56. Natural levce deposits between Palazuelo de Boñar and La Ercina. Asymmetrical smallscale channel (A) filled with sand and organic matter (dark coloured) indicates current direction.

Fig. 57. Exposure S of La Mata de la Riba. The letters refer to the subdivision made in Fig. 15.

Fig. 58. Differentiation between quartz grains and feldspar grains by means of hemateine and cobaltinitrite. Quartz: dark coloured and glass-like (Q); feldspars: light coloured and not translucent (F). Sample Va2.





#### PLATE VII

Fig. 59. Rolled-up vermicular aggregate of kaolinite (K) of detrital origin; grain shape is accentuated by a thin iron skin. Sample 326, 100x.

Fig. 60. Muscovite flake (M), strongly compressed, crumpled and split, of detrital origin. Sample 049, 100x

Fig. 61. Rutile skeletons (sagenite) in discoloured biotite grain (B). Sample 043, 100x.

Fig. 62. Biotite grain (B) which escaped alteration as a result of embedding in a quartz fragment of detrital origin (Q1) and closing off by authigenic quartz (Q2). Sample 001, 40x.

Fig. 63. Quartz grain (Q) strongly corroded by solution. Sample 326, 100x.

Fig. 64. Kaolinite sheaves (K) formed as cement between detrital quartz grains. Sample 001, 100x.

Fig. 65. Kaolinite sheaf (K), strongly compressed between detrital quartz grains (Q). Sample 306, 100x.

Fig. 66. Kaolinitic cement (K) partly replaced by iron (hydr)oxides (IO) with formation of radial aggregates (R) in the transition zone. Sample 032A, 25x.





#### PLATE VIII

Fig. 67. Stringers consisting of black organic matter showing a parallel orientation perpendicular to the plane of stratification. Sample 070, 25x.

Fig. 68. Paleozoic sandstone showing good sorting and concavo-convex and sutured contacts between the detrital grains. Dust-rings are common. Sample 271, 40x.

Fig. 69. (cf. Fig. 68) Sample of Voznuevo sandstone showing poor sorting and point and straight contacts between the detrital grains. Dust-rings are rare. Note abundance of cement. Sample 090, 100x.

Fig. 70. Radiolarite of the Vegamián Formation. Sample 112, 40x.

Fig. 71. Radiolarite of the Voznuevo Formation. Sample 021, 40x.

Fig. 72. Strongly folded phtanite, probably of Precambrian origin. Sample 132, 40x.

Fig. 73. Strongly folded quartz phyllite, probably of Precambrian origin. Sample 093C, 25x.

Fig. 74. Cellular structure of wood fibre (fusite) in charcoal-like fragment. Sample 249, 100x.











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Rob. K. JONKER. Leidse Geologische Mededelingen, deel 48. ENCLOSURE III

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### Section J

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ರೆರೆ		L = 4 cm clayey silt silt col: wh clay col: gy L = 4 cm	80	C 21	800 800		day L= 30 cm L= 4 cm Iy: silt with coarse sand				
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![](_page_26_Figure_1.jpeg)

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